



Validating improved-MODIS products from spectral mixture-Landsat snow cover maps in a mountain region in southern Spain

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Abstract. Remote sensing is the only feasible data source for distributed modelling of snow in mountain regions on medium to large scales, due to the limited access to these areas together with the lack of dense ground monitoring stations for snow variables. Observations worldwide identify snow cover persistence together with snowfall occurrence as the most affected variables by global warming. In Mediterranean regions, the spatiotemporal evolution of the snow cover can experiment quick changes that result in different accumulation-ablation cycles during the cold season. High frequency sensors are required to adequately monitor such shifts; however, for trend analyses, the Landsat time series constitute the only available source of data, being their frequency low for this regime, especially when cloudy conditions limit the available images. On the other hand, the MODIS daily series provide more than 15 years of continuous snow maps, despite the spatial resolution may pose a constraint in areas with abrupt topography; several approaches have been done to improve their spatial resolution from combining different information. This work presents a methodological approach to validate the improved MODIS daily snow cover maps from Notarnicola et al. (2013a, b), with 250 m spatial resolution, in Sierra Nevada (southern Spain), from a reference data set obtained by spectral mixture analyses of Landsat TM data by Pimentel et al. (2017b). This reference time series of fractional snow maps, with 30 m spatial resolution, were validated from high resolution local time series of snow maps obtained by terrestrial time-lapse cameras. The results show a significantly high correlation between the two snow map products both on a global and basin scales in the Sierra Nevada area. Selected areas and time periods are shown to address the convergence and divergence between both products and assess the development of a fusion algorithm to retrieve daily Landsat-resolution snow maps on a long term basis.

1 Introduction

Remote sensing techniques constitute the best source to provide distributed information about the snowpack evolution on medium to long time scales, complementing the traditional in situ field surveys and automatic ground measurements. Snow cover fraction (SCF) is one of the more reliable snow related variable measured from the space (Dozier and

Painter, 2004) and is commonly used in hydrological studies to calibrate, evaluate, or be assimilated into snow distributed modelling (Andreadis and Lettenmaier, 2006; Parajka and Blöschl, 2008; Pimentel et al., 2015). Within the different Earth Observation (EO) missions, (1) Landsat-5 (TM), Landsat-7 (ETM+) and Landsat-8 (OLI), with 30 × 30 m spatial resolution and 16 days revisiting time (Roy et al., 2014; Pimentel et al., 2017a), and (2) MODIS Terra and Aqua, with

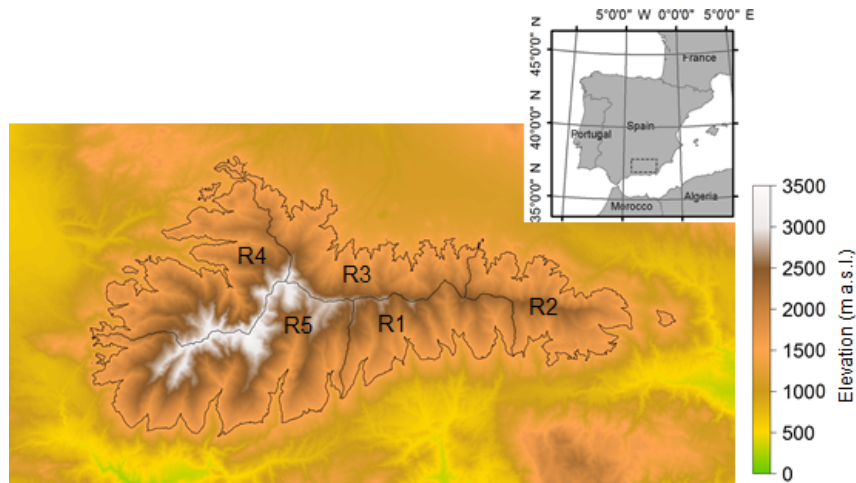


Figure 1. Area of Sierra Nevada Mountain (Spain) above 1500 m a.s.l. and limits of the five regions in which the study area has been divided for the spatial analysis: R1 – Adra, R2 – Andarax, R3 – Fardes, R4 – Genil and R5 – Guadalfeo.

500 m grid cell resolution and daily temporal frequency for snow, are the most extended data sources for snow studies, since they offer the highest spatial and temporal resolution, respectively (Hall et al., 2002).

However, Mediterranean mountainous areas are extremely vulnerable to climate change effects and highly dependent on snow water resources (Barnett et al., 2005; Giorgi, 2006). The particularities of the snowpack make the use of raw EO products not enough to capture these specific patterns. For instance, the very strong spatiotemporal variability, which very often undergoes different accumulation-snowmelt cycles during the cold season in a given year, or the snow patched distribution around local singularities, such as rocks and vegetation, consequence of a very complex ablation process (Ménard et al., 2014; Pimentel et al., 2015, 2017b). Hence, both high temporal and spatial resolutions are required to have a realistic representation of the snow cover.

In this context Pimentel et al. (2017a) carried out a spectral mixture analysis to derive fractional snow cover map time series from Landsat TM and ETM+. High resolution terrestrial photography (TP) was used as ground truth to validate the obtained product. The spatial resolution of the snow cover area was improved using this technique; however, the large revisiting time of Landsat TM and ETM+ in addition to the presence of clouds in some of the dates with snow presence, constitute a big constraint for useful time series. Using the same idea of improving spatial representation of the snow, Notarnicola et al. (2013a, b) developed an algorithm that combines different MODIS products: MOD09GQ-MYD09GQ, MOD09GA-MYD09GA, MOD021KM-MYD021KM, MOD03-MYD03, to produce snow cover maps at 250 m spatial resolution and daily frequency. The algorithm has specific modules to take into account the effect of vegetation and clouds and increase

the spatial resolution of the standard snow MODIS product, MOD10A1-MOD10A2, from 500 to 250 m.

This work presents a methodological approach to assess the improved MODIS daily snow cover maps from Notarnicola et al. (2013a, b), in Sierra Nevada (southern Spain), using as reference data set the Landsat fractional cover maps obtained by spectral mixture analyses by Pimentel et al. (2017a).

2 Study site and data available

This study is carried out in Sierra Nevada Mountains, southern Spain (Fig. 1). They are a linear mountain range of 90-km length that runs parallel to the coastline of Mediterranean Sea. Alpine and Mediterranean climate conditions coexist in just a 40-km distance. Strong altitudinal gradients with marked differences between the south (directly affected to the sea) and the north faces are found in the area.

The snow usually appears above 2000 m a.s.l. during winter and spring even though the major snowmelt season generally lasts from April to June, but can be also found at lower altitudes every year. The typically mild Mediterranean winters produce several snowmelt cycles before the final melting phase, which distributes the snow in patches over the terrain. Precipitation and temperature regimes are highly variable among years, with annual precipitation values averaged in the area that can range from 200 to 900 mm and annual mean of the daily minimum and maximum temperature of -5 and 30 °C, respectively (Pérez-Palazón et al., 2015).

Two snow cover EO products are used in this study: (1) Fractional snow cover maps, at 30×30 m and 16 days spatial and temporal resolution respectively, derived from spectral mixture analysis of Landsat TM and ETM+ validated using as high resolution terrestrial photography (Pimentel et al., 2017a); (2) binary snow cover maps obtained

from an algorithm that combines several MODIS products and produce daily snow cover maps with a grid cell size of 250 m (Notarnicola et al., 2013a, b). In the text these products will be referred as Landsat-mix and MODIS-EURAC, respectively. Both products have a temporal overlapping from 1 January 2002 to 31 August 2013. For this period, a total number of 108 and 2963 cloud-free images were used in the study for Landsat-mix and MODIS-EURAC, respectively; considering as cloudy images those whose presence of clouds exceeded a 10 % of the study area.

3 Methods

SCF was calculated over the whole study area in Sierra Nevada (area above 1500 m a.s.l) and in each one of the five main headwaters regions: R1 – Adra, R2 – Andarax, R3 – Fardes, R4 – Genil and R5 – Guadalfeo, for both snow products, Landsat-mix and MODIS-EURAC.

SCF from both products were compared in the 108 common dates. A simple linear model was fitted in those days to relate both products. Landsat-mix was chosen as ground truth and MODIS-EURAC as dependent variable. Equation (1) was used for that,

$$SCF_{\text{Landsat-mix}} = (SCF_{\text{MODIS-EURAC}} \times a) + b, \quad (1)$$

where a and b are the two parameters of the lineal model. Using this model, the SCF from MODIS-EURAC was reconstructed using the 2963 cloud-free images.

4 Results

Figure 2 shows the evolution of SCF from MODIS-EURAC and Landsat-mix in the overlapping dates for both products in each of the defined regions and over the whole study area. SCF follows the same trend for both products, with a clear overestimation of the SCF derived from MODIS-EURAC. This overestimation is especially significant during the dates with higher SCF values and specifically in R1 – Adra, where differences about $0.20 \text{ m}^2 \text{ m}^{-2}$ can be found in the 2 days with more snow throughout the study period. Differences are practically negligible during dates with low SCF for all the regions. However, this general overestimation trend from MODIS-EURAC change during the last stages of the snowmelt season, when its lower resolution is not able to capture snow remaining isolated snow patches.

Figure 3 and Table 1 show the linear relation found between the two products and the parameters that fitted these relationships respectively. The linear pattern is clear for all regions, with determination coefficients ranging from 0.979 for R4 – Genil to 0.995 for R2 – Andarax, with a clear overestimation of MODIS-EURAC.

The parameter a (Table 1), which measures the slope of the fitted model and consequently determines the magnitude

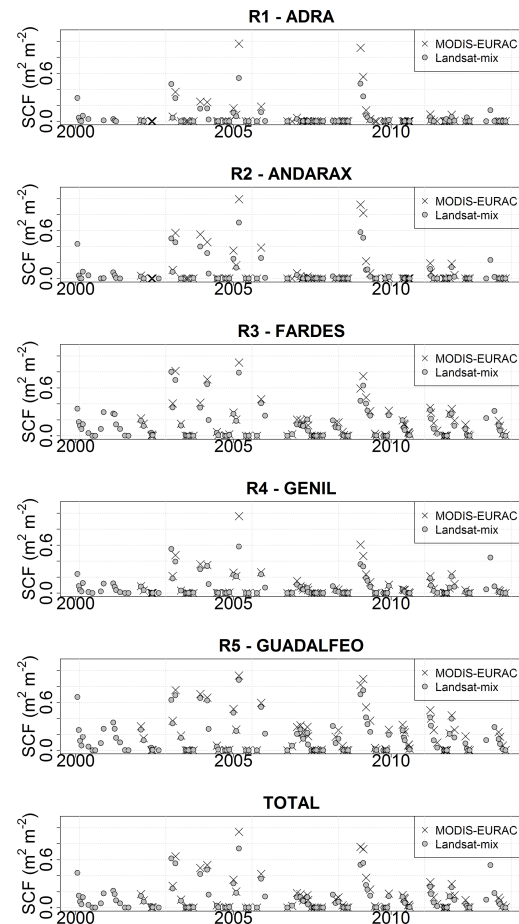


Figure 2. Comparison between SCF evolution of both products MODIS-EURAC (black crosses) and Landsat-mix (gray dots) in the 108 overlapping dates, in each of the regions selected and in the whole study area.

of the MODIS-EURAC overestimation, differs between regions. Lower values of the parameters, which imply higher overestimations, are found in the drier and warmer areas (Pérez-Palazón et al., 2015), 0.700 in R1 – Adra and 0.645 in R2 – Andarax; located in the south face and with lower mean elevation. On the contrary, wetter and colder regions have higher values and consequently less overestimation coming from MODIS-EURAC. Although, the general accuracy from MODIS snow products is estimated approximately at 93 % (Hall and Riggs, 2007) and similar studies has found an accuracy of 94.6 % comparing MODIS products with surface observations over northern China, (Huang et al., 2016), the heterogeneity of the snow distribution due to the abrupt terrain and climate conditions, make the overestimations of MODIS-EURAC over this are slightly bigger.

The clear linear fit found allows using this relationship as a simple model to correct the average values calculated using MODIS-EURAC over the study area. Overestimation corrections of 0.30, 0.35, 0.13, 0.20 and 0.23 $\text{m}^2 \text{ m}^{-2}$, were achieved for Adra, Andarax, Fardes, Genil and Guadalfeo,

Table 3. Annual maximum average snow cover fraction (SCF) calculated for the three snow products, Landsat-mix (Land), MODIS-EURAC (MOD) and Corrected MODIS-EURAC (NEW) in each of the study regions and the whole study area.

	Region 1			Region 2			Region 3			Region 4			Region 5			Total		
	Land	MOD	NEW	Land	MOD	NEW	Land	MOD	NEW	Land	MOD	NEW	Land	MOD	NEW	Land	MOD	NEW
2002–2003	0.011	0.306	0.214	0.024	0.419	0.270	0.188	0.549	0.472	0.082	0.457	0.362	0.259	0.643	0.496	0.146	0.449	0.348
2003–2004	0.467	0.745	0.521	0.501	0.639	0.412	0.800	0.969	0.833	0.552	0.791	0.626	0.690	0.781	0.602	0.616	0.732	0.567
2004–2005	0.160	0.734	0.513	0.400	0.853	0.550	0.647	0.672	0.577	0.338	0.656	0.519	0.654	0.840	0.648	0.474	0.742	0.575
2005–2006	0.541	0.955	0.668	0.701	0.965	0.622	0.790	0.803	0.690	0.582	0.972	0.770	0.883	0.878	0.677	0.739	0.893	0.692
2006–2007	0.115	0.864	0.605	0.257	0.737	0.475	0.408	0.917	0.789	0.234	0.694	0.549	0.542	0.884	0.681	0.360	0.773	0.599
2007–2008	0.027	0.498	0.348	0.043	0.556	0.359	0.204	0.757	0.651	0.106	0.879	0.696	0.261	0.897	0.691	0.131	0.689	0.534
2008–2009	0.006	0.472	0.330	0.027	0.353	0.228	0.193	0.700	0.602	0.086	0.382	0.303	0.304	0.572	0.441	0.162	0.368	0.285
2009–2010	0.471	0.857	0.599	0.579	0.938	0.604	0.628	0.911	0.784	0.362	0.628	0.497	0.750	0.908	0.700	0.558	0.778	0.603
2010–2011	0.015	0.986	0.690	0.021	0.895	0.577	0.256	0.877	0.754	0.091	0.981	0.777	0.249	0.883	0.681	0.142	0.831	0.644
2011–2012	0.051	0.421	0.294	0.119	0.581	0.374	0.323	0.832	0.716	0.183	0.563	0.446	0.411	0.851	0.656	0.262	0.659	0.511
2012–2013	0.054	0.552	0.387	0.144	0.559	0.360	0.279	0.930	0.800	0.206	0.917	0.727	0.396	0.823	0.634	0.254	0.701	0.543

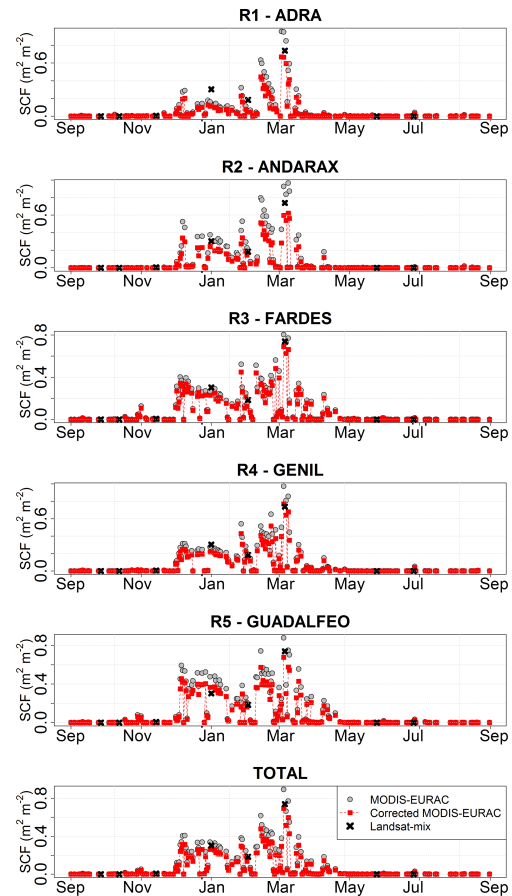


Figure 4. Example of reconstruction of MODIS-EURAC snow cover map for the year 2004–2005.

EURAC-MODIS product presented a general overestimation during the high snow covered period, which was partially solved with the correction with the linear model. Further on-going work is exploring the apparent threshold in the large SCF values domain that can be observed in the graphs, together with the different behaviour of Region 5, the mostly-influenced by snow in the study area.

5 Conclusions

This work shows how the setting out of a simple approach provides a more accurate evolution of the average SCF values over this Mediterranean region, combining the advantages of two already existing products, the high spatial accuracy of Landsat-mix and the daily temporal resolution of MODIS-EURAC. The result is a daily time series on which different studies that require high resolution both of time and space can be based on. This work constitutes the first step in a more complex development of a data fusion algorithm that not only reproduces average behaviour but also snow distribution at grid/subgrid scales.

Data availability. The improved-MODIS snow products used in this study are available at <http://sdi.eurac.edu/geonetwork/srv/eng/metadata.show?id=357955&currTab=advanced> (Notarnicola et al., 2013a, b).

Author contributions. RP has been the main responsible of the calculation and writing; CM, LDG and CN have processed the improved-MODIS product; MJPP has contribute in the Landsat scene processing; and MJP has provided the initial idea of this study and help in the writing. All the authors have contributed in the review process.

Competing interests. The authors declare that they have no conflict of interest.

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